# MESOZOIC CALCAREOUS NANNOFOSSILS FROM MASIRAH ISLAND (SULTANATE OF OMAN)

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Abstract:Calcareous nannofossils were studied in support of the geological mapping and structural interpretation of Masirah Island in Oman. The assemblages found include Middle Triassic(?), Early Jurassic (Late Pliensbachian), and Early as well as Late Cretaceous assemblages from a little-known region between the western Tethys and the Pacific Realm.

#### Introduction

Masirah Island is one of the less well-known fragments of ocean floor occurring at the southeastern coast of Oman (Figure 1). The island extends over roughly 700km² and is formed by two well-exposed ophiolite nappes (Marquer *et al.*, 1995) presently under study by a group of earth scientists at the University of Berne and Neuchatel (Peters *et al.*, 1995).

crust of the evolving proto-Indian Ocean which opened in the Late Jurassic between West and East Gondwanaland (e.g. Salman & Abdulah, 1995). Masirah became part of the northward drifting Indian Transit Plate (including Pakistan and the Seychelles), when this plate separated from Africa/Madagascar after 84Ma (Immenhauser et al., 1997). The ophiolite was obducted in its present position onto the continental margin of SE Arabia at Cretaceous/Terti-

MASIRAH ISLAND Gulf o ARABIA Oman oniacian/Santonian-Lov I-257C Coniacian/Sant I-260C Coniacian/Santonian-I-261C Coniacian/Santo AFRICA I-188C Turonian South of Umm Ar Rusays 1-356C Uppermost Maastrichtian Wadi Al Firay I-319C Tithonian-Berria Wadi Thoumi I-210C Tithonian-Lower Aptia Wadi Maghilah I-121C Berriasian-Hauterivian Log XXI I-329C Berriasian-Valanginian I-330C Berriasian-Hauterivian I-331C Berriasian-Valanginian I-332C Berriasian-Valanginian /Valanginian-H I-643C Campanian Maastrichtian Fayah Unit I-335C Tithonian/Berriasian Wadi Haqi I-515 Conlack I-281C Albian-33R Aptian or Lower Albian I-637R Mid-Upper Albian North of Suwayr Kalban Unit Maghilah Unit Ophiolite/ I-290C Albian/Cenomai S. of Jabel Suwayr Tertiary/Ouaternary

Figure 1: Map of Masirah Island, Sultanate of Oman, showing the localities yielding useful calcareous nannofossil assemblages.

The Masirah Ophiolite was formed about 150Ma (Figure 2) during the Late Jurassic/Early Cretaceous at a latitude of 38°±12°S (Gnos & Perrin, 1996; Immenhauser, 1995, 1996). It belongs to the oldest-preserved oceanic

ary boundary times, around 65Ma. The obduction is thought to be related to oblique, hot-spot (Deccan Traps)-induced rifting between India/Pakistan and the Seychelles Microcontinent at 65Ma (Plummer & Belle, 1995; Immenhauser *et al.*, 1997).

The autochthonous sedimentary cover of the lower of the two ophiolite nappes on Masirah provides an unbroken stratigraphic record from 150Ma to 65Ma (Immenhauser, 1995). The dating of samples of the highly tectonised sediments was achieved by calcareous nannofossils, foraminifera and radiolaria (Dumitrica et al., 1997). Much to the surprise of the authors, also the Middle Triassic yielded questionable calcareous nannofossils, and Lower Jurassic calcareous nannofossil assemblages were found besides the expected Lower and Upper Cretaceous ones. Since the knowledge about Triassic and Lower Jurassic calcareous nannofossils from the southeastern part of the Tethys is small, we decided to present the new findings in order to complement the present palaeogeographic distribution pic-

Since reports about calcareous nannofossils from other time periods of this part of the world are scarce too, we also present the results of the nannofossil analyses from later time periods and discuss the problems encountered in the age assignment of the various sediment packages.

### Material and methods

Spot-samples for age determination by coccoliths were collected from pelagic sediments and were mainly taken in the shaley partings and interbeds of radiolarian micrites

Figure 2: Palaeotectonic map of the Toarcian, modified after Ricou (1994), with position of Masirah Island and Lower Jurassic calcareous nannofossil occurrences after de Kaenel & Bergen (1993), Bown (1987), Erba et al. (1992) and Kristan-Tollmann (1988). Triassic occurrence after Bralower et al. (1991).

and even from ribbon-bedded cherts. Upper Cretaceous flysch-type sediments were best dated with samples taken in shaley interbeds. Simple smear-slides were prepared from the samples and embedded in Canada balsam. The study was undertaken at a magnification of x1000 with a light microscope (LM) equipped with polarisers and a gypsum plate. Approximately 250 smear-slides were prepared of which some 20% yielded more or less datable coccolith floras. A few samples were also studied by scanning electron microscope (SEM), the preparation following the usual techniques as described in Perch-Nielsen (1985).

#### **Middle Triassic**

Sample I-620, which yielded Middle Triassic radiolaria also included forms which - with doubt - may possibly be assigned to the calcareous nannofossil genus Calyculus. This genus was described by Noël (1973) from the Toarcian (Lower Jurassic) of France and consists of subvertical elements surrounding a relatively wide central area. According to Erba et al. (1992), "Calyculus has not been reported from older than Toarcian sediments so far". While the organic origin of the specimens illustrated in Plate 1, Figures 1-3 seems likely, their generic assignment must remain speculative. No other calcareous nannofossils were found by SEM nor LM in sample I-620, which yielded the following age-indicative radiolaria: Tritortis balatonica Kozur and Tritortis kretaensis subcylindrica Kozur (Upper Langobardian or Lower Cordevolian; Dumitrica et al., 1997).

## Lower Jurassic

The age assignment of Early Jurassic was essential in proving the presence of exotic sediment blocks older than the formational age of the ophiolite on which they were found. Exotic sediment blocks on Masirah Island form part of the Upper Cretaceous Fayah Unit and were shed in the latest Maastrichtian (see below).

Three samples (I-240, I-243 and I-244; Figure 3) were found to include the very rare to rare, poorly to moderately well-preserved calcareous nannofossils listed in Figure 4. Sample I-243 was only studied by SEM. The illustrations are shown on Plate 1. The fact that the assemblages are different is probably due to the different technique used for their study in the case of I-240 and I-243, and by the very poor state of preservation of the coccoliths

Figure 3: Distribution of Upper Pliensbachian calcareous nannofossils from Masirah Island																		
SAMPLE	ABUNDANCE	PRESERVATION	Biscutum profundum	Calcivascularis jansae	Calyculus ? sp.	Crepidolithus crassus	Crepidolithus granulatus	Lotharingius sp.	Mitrolithus elegans	Mitrolithus lenticularis	Mitrolithus sp.	Orthogonoides hamiltoniae	Parhabdolithus robustus	Schizosphaerella punctulata	Similiscutum avitum	Similiscutum crucíulus	Tubirhabdus patulus	Zeugrhabdotus erectus
240	R	Р	+	+	+	+	+	+	+	+	+	+	+	+			+	+
243	vR	Р	+												+	+		-
244	vR	vΡ	+	+			+							+				

in sample I-244. For LM illustrations see Crux (1987) and de Kaenel & Bergen (1993), who also present SEM pictures.

#### Age

Only sample I-240 can be used for an age determination. The presence of *Lotharingius* sp., the first, small form of that genus, suggests an age not older than late Early Pliensbachian (*davoei* Zone of Northern Europe and Portugal, according to de Kaenel *et al.* (1996)). The presence of *Biscutum profundum* means, however, that the sample is of even younger, Late Pliensbachian age (*margaritus* Zone). At the same time, the sample cannot be younger, since it includes *Parhabdolithus robustus*, the last occurrence (LO) of which has been reported in the same zone by de Kaenel *et al.* (1996). This age assignment is contradicted by the distributions reported by Bown (1996) who reported the LO of *P. robustus* in the Lower Pliensbachian *ibex* Zone.

## Palaeogeography

The assemblage in sample I-240 is as rich as coeval assemblages from the western and eastern Tethys or from Northern Europe, and no forms restricted to Oman were found. From the palaeogeographic map (Figure 2), it follows that the Lower Jurassic coccolith assemblage reported here was deposited about mid-way between the Italian Jurassic assemblages which were classified as 'tethyan' and those labelled 'pacific' in Timor by Bown (1987). While one Masirah assemblage includes Mitrolithus jansae, which was found by Bown (1987) to be restricted to an equatorial/sub-equatorial belt and absent in the Pacific samples from Argentina and the Queen Charlotte Islands (Bown, 1992), none includes Mazaganella pulla or Mazaganella protensa that were reported by the same author from the southern edge of the Tethys. The sample also lacks Crucirhabdulus primulus and Parhabdolithus liasicus, two species usually found in Pliensbachian assemblages.

#### Cretaceous

A synthetic composite section of the autochthonous Cretaceous sediments on the lower nappe of the Masirah Ophiolite, as reconstructed from observations in the field and the ages derived from coccolith, foraminiferal and radiolarian dating, can be found in Immenhauser (1995).

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Four groups of Cretaceous assemblages could easily be distinguished on the basis of their calcareous nannofossil content (Figures 4, 5). They are dated, where possible, according to the zonation proposed by Sissingh (1977), taking into account also the findings of Perch-Nielsen (1985) and Bralower *et al.* (1989).

The sediments that were assigned an Early Cretaceous age include only poorly- to very poorly-preserved

coccoliths. Their distribution in some samples is presented in Figure 4. The age assignment of many samples was difficult and often represents an educated guess rather than a firm statement. Nevertheless, the presence of forms generally used for the biostratigraphic subdivision of the Lower Cretaceous of the Tethyan Realm allows for an approximate age determination of a dozen or so samples. The knowledge of the presence of these species in this part of

the Tethys may prove useful in future studies. Classical markers used in the Tethyan Realm for zonation found herein are: Calcicalathina sp.(?), Cruciellipsis cuvillieri, Eprolithus floralis, Microstaurus chiastius, Polycostella beckmannii, P. senaria(?), Prediscosphaera columnata, Rucinolithus irregularis and various species of Nannoconus. Forms known mainly from the Boreal Realm include Kokia sp.(?) and Lithraphidites morayfirthensis(?). The poor preservation does not allow for an unequivocal determination of the forms indicated by '(?)'.

Three major Upper Cretaceous assemblages were found (Figure 5): the Coniacian/Santonian is characterised by the presence of Marthasterites furcatus (which, however, ranges from the Late Turonian through the earliest Campanian, but is most common in the Coniacian and Santonian). The Campanian is approximated by the presence of Broinsonia parca parca and the Upper Maastrichtian by the presence of Micula murus, the marker-species for the uppermost Maastrichtian. Other Upper Cretaceous marker-species include Arkhangelskiella cymbiformis, Eiffellithus eximius, Helicolithus trabeculatus, Lithastrinus septenarius, Lithraphidites quadratus, Prediscosphaera stoveri, Reinhardtites anthophorus and Tranolithus orionatus.

Upper Cretaceous sediments from Oman were the object of a special study on the evolutionary trends in the family Arkhangelskiellaceae by Lauer (1974). In our material, specimens of this family were not common and thus were not restudied especially. An intensive search in the field for samples above those that include Micula murus, in order to find sediments including Micula prinsii, the marker of the terminal Cretaceous, or even a Cretaceous/Tertiary boundary section, was not successful. The contact with the se diments overlying those including M. murus is tectonic.

## Tertiary

Very rare Tertiary coccoliths were found in several samples. They include Coccolithus pelagicus and small Prinsiaceae, such as are found from the Paleocene through the Recent, and Chiasmolithus solitus (Eocene). It has not been possible to ascertain that these occurrences are actually dating the sample in question. Since many samples included no coccoliths at all, contamination in the laboratory seems, however, unlikely.

#### Systematic remarks

Systematic information and illustrations of the calcareous nannofossil species listed, mentioned and illustrated can be found in Perch-Nielsen (1985), Bown (1987) and de Kaenel & Bergen (1993).

#### Discussion and conclusion

What had started out as a routine biostratigraphic investigation of samples collected by a geologist in the field in order to help him with the interpretation of the complicated geology, turned out to actually change the previous

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geological picture considerably. In addition, it provided an interesting surprise in the form of enigmatic Middle Triassic nannofossils and an Early Jurassic calcareous nannofossil assemblage from an area from which none was previously described. This assemblage could be shown to correspond broadly to those found in the western part of the Tethys.

According to Bown (1987), the relatively common presence of *Calcivascularis jansae* "reveals the tethyan affinities of a nannofossil assemblage immediately". At the same time, it seems that some of the species usually present in the tethyan Pliensbachian, such as *Crucirhabdus primulus* and *Parhabdolithus liasicus*, are missing in the samples studied.

The often relatively poor state of preservation of the Lower Cretaceous calcareous nannofossil assemblages prevented the assignment of most samples to a specific zone. Nevertheless, a more broad age assignment was possible in many cases and sometimes supported by radiolarian stratigraphy. Also, the Upper Cretaceous assemblages were rarely well enough preserved to allow for a clear zonal assignment.

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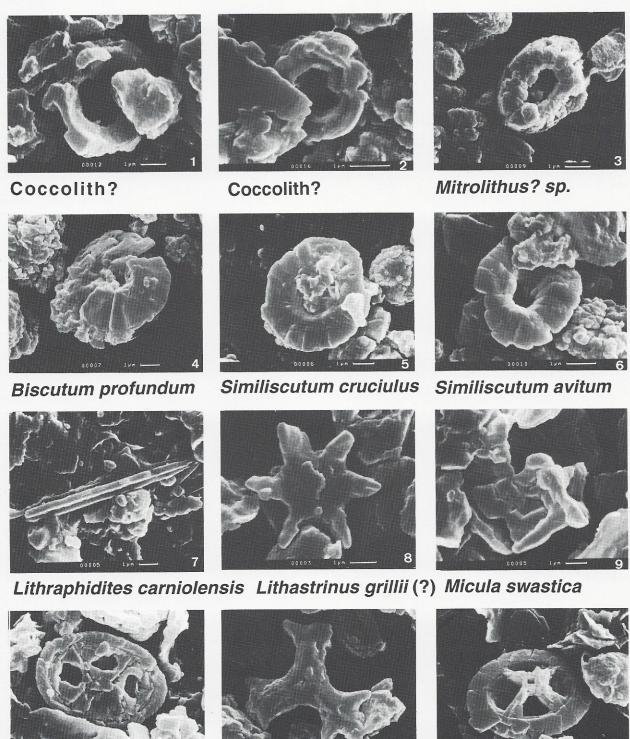
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#### PLATE 1

- **Figs 1, 2: Coccoliths or inorganic calcite grains?** Sample I-620 of Middle Triassic age based on radiolaria (Dumitrica *et al.*, 1997). Magnifications: x7000 & x11350.
- Fig.3: Mitrolithus? sp. Sample I-243, Pliensbachian, x7000.
- Fig.4: Biscutum profundum de Kaenel & Bergen (1993). Sample I-241, Pliensbachian, x5000.
- Fig.5: Similiscutum cruciulus de Kaenel & Bergen (1993). Sample I-241, Pliensbachian, x6000.
- Fig.6: Similiscutum avitum de Kaenel & Bergen (1993). Sample I-243, Pliensbachian, x8000.
- Fig.7: Lithraphidites carniolensis Deflandre (1963). Sample I-356, Upper Maastrichtian, x5000.
- Fig.8: Lithastrinus grillii? Stradner (1962). Sample I-257, Coniacian/Santonian, x7600.
- Fig.9: Micula swastica Stradner & Steinmetz (1984). Sample I-356, Upper Maastrichtian, x6300.
- Fig.10: Helicolithus trabeculatus (Górka, 1957) Verbeek (1977). Sample I-261, ±Coniacian/Santonian, x6000.
- **Fig.11:** *Marthasterites furcatus* (Deflandre *in* Deflandre & Fert, 1954) Deflandre 1959. Sample I-257, ±Coniacian/Santonian, x6250.
- Fig.12: Prediscosphaera cretacea (Arkhangelsky, 1912) Gartner (1968). Sample I-261, ±Coniacian/Santonian, x6500.

## PLATE 1



Helicolithus trabeculatus Marthasterites furcatus Prediscosphaera cretacea